

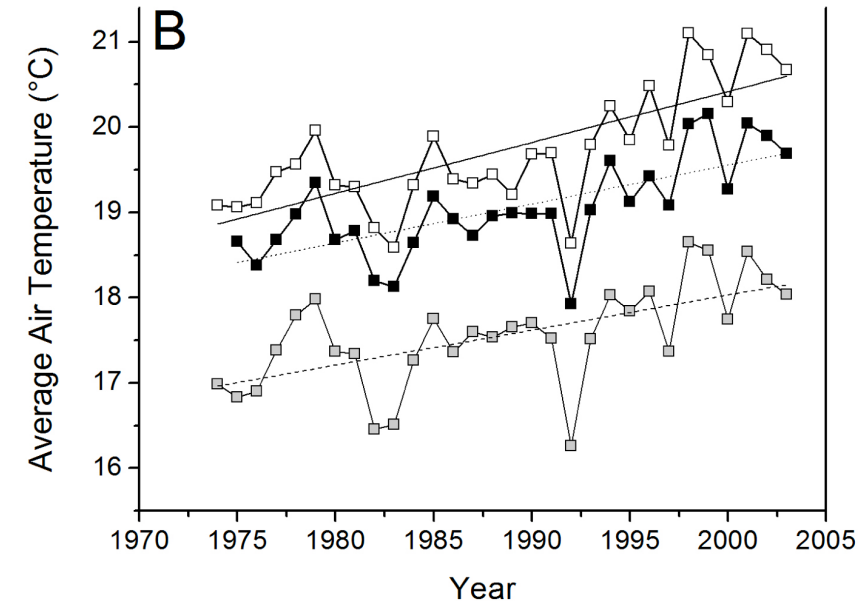
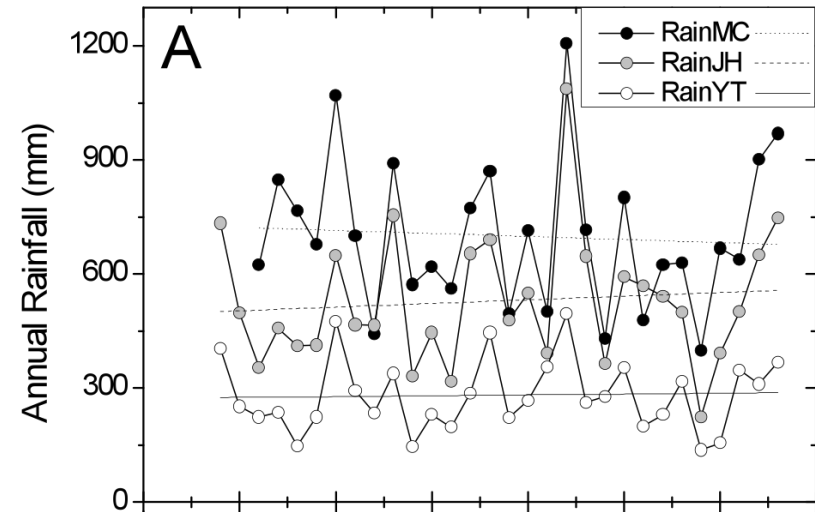
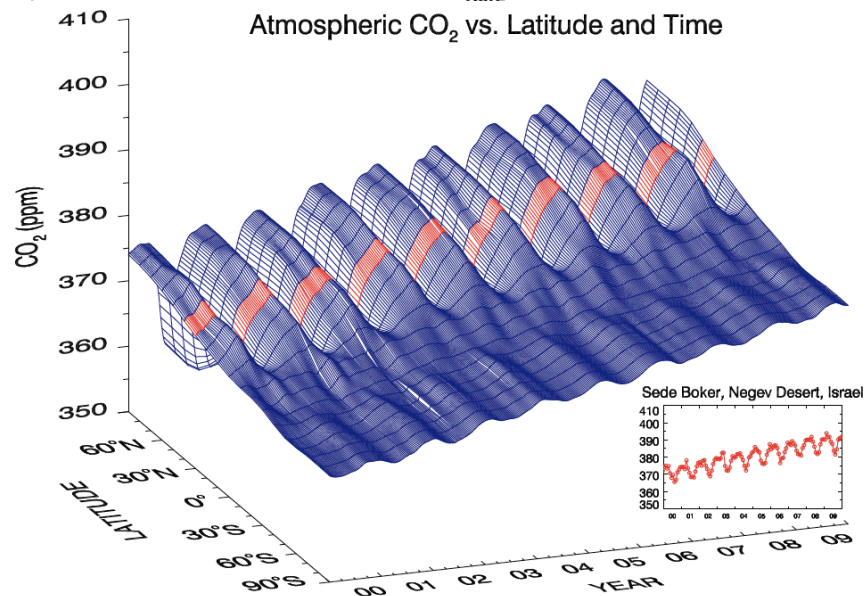
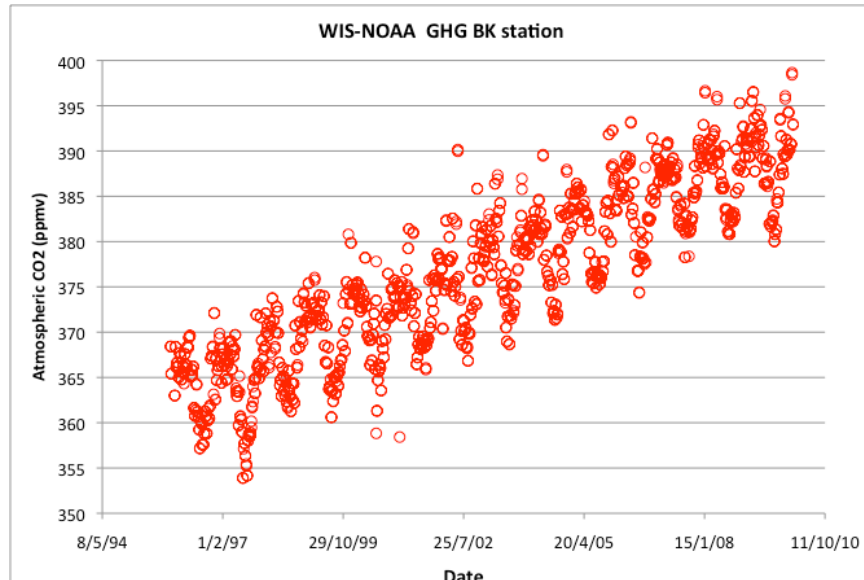


Tradeoffs in semi-arid pine forest

KKL; GLOW-JR; ISF

Dan Yakir & Eyal Rotenberg

ISEEQS, June 2010



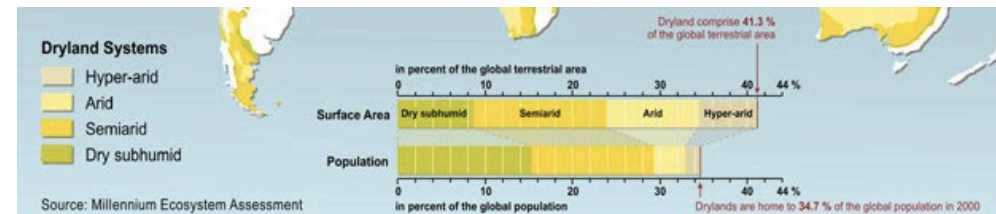
The global FluxNet



Extent of Drylands

Region	Arid		Semi		Dry Sub-Humid		All Drylands		Population
		%		%		%		%	
Asia (Incl. Russia)	6,164	13	7,649	16	4,588	9	18,401	39	1,444,906
Africa	5,204	17	5,073	17	2,808	9	12,933	43	267,563
Oceania	3,488	39	3,532	39	996	11	8,016	89	140,586
North America	379	2	3,436	16	2,081	10	5,896	28	86,990
South America	401	2	2,980	17	2,233	13	5,614	32	59,323
C. America & Caribbean	421	18	696	30	242	10	1,359	58	31,719
Europe	5	0	373	7	961	17	1,339	24	6,960
	15,910	12	23,739	18	13,909	10	53,558	41	2,038,047

Land area: millions of square kilometers. Population: thousands. From White and Nackoney (2003)





Indicators of high carbon use efficiency



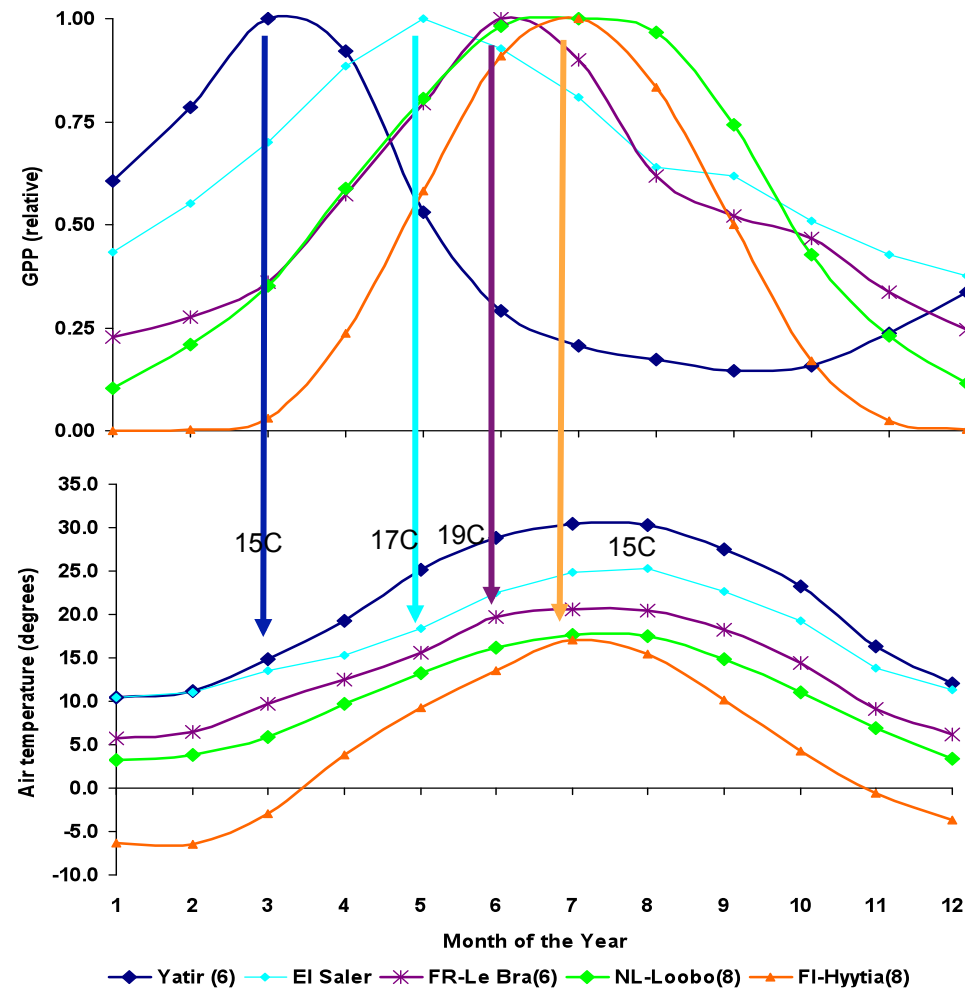
Pine forest	GPP	Re	NEE	NEE/GPP
European (Carboeurope)	1142	944	200	0.17
Global (FluxNet)	1540	1280	260	0.17
Semi-arid (Yatir)	820	600	220	0.27

*gr C m⁻²

Maseyk et al., 2007a,b; Grunzweig et al., 2009

Large-scale “homeostasis” response

$$\text{DOY}_{\text{GPP}} = -0.71(R_G) + 263.1; R^2 = 0.87$$

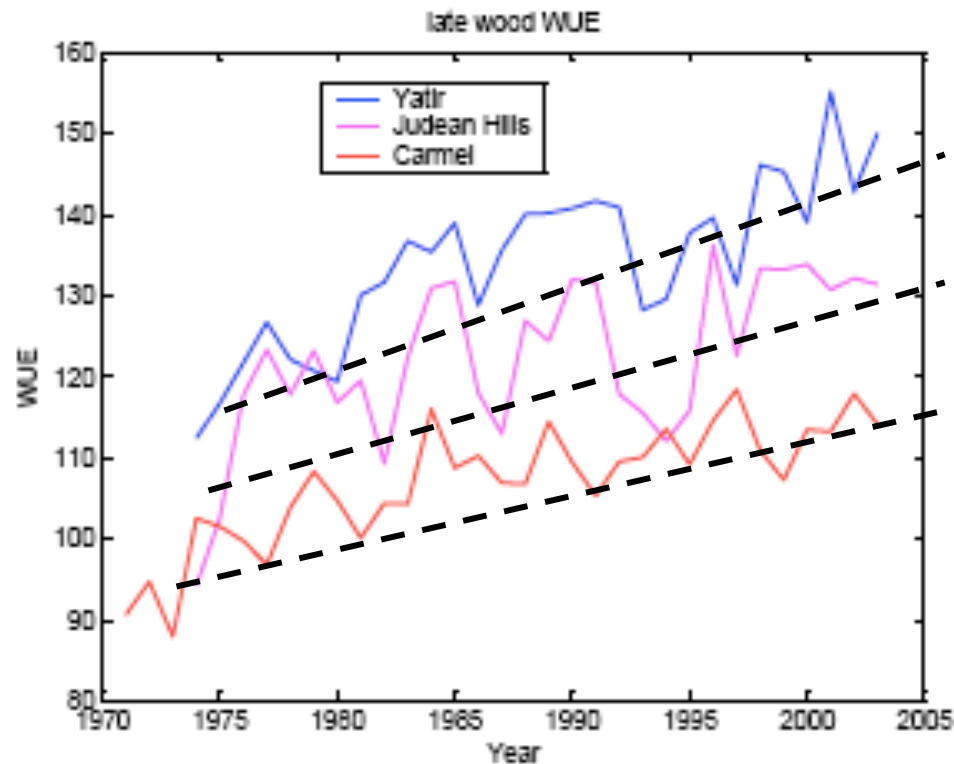


Rotenberg & Yakir, Science 2010

Annual pattern of a.. the monthly relative gross primary productions (defined as ((monthly sum of GPP)/(sum of highest productive GPP month)), and b.

^{13}C derived WUE along precipitation gradient

$\sim 25\%$ increase in WUE, in dry sites: CO_2 response...



CO_2 up (330 to 380 ppm)
 T up $\sim 1.5^\circ\text{C}$, no IA correlation
 P no change, strong IA correlation



Episodic events vs. long-term climate change...

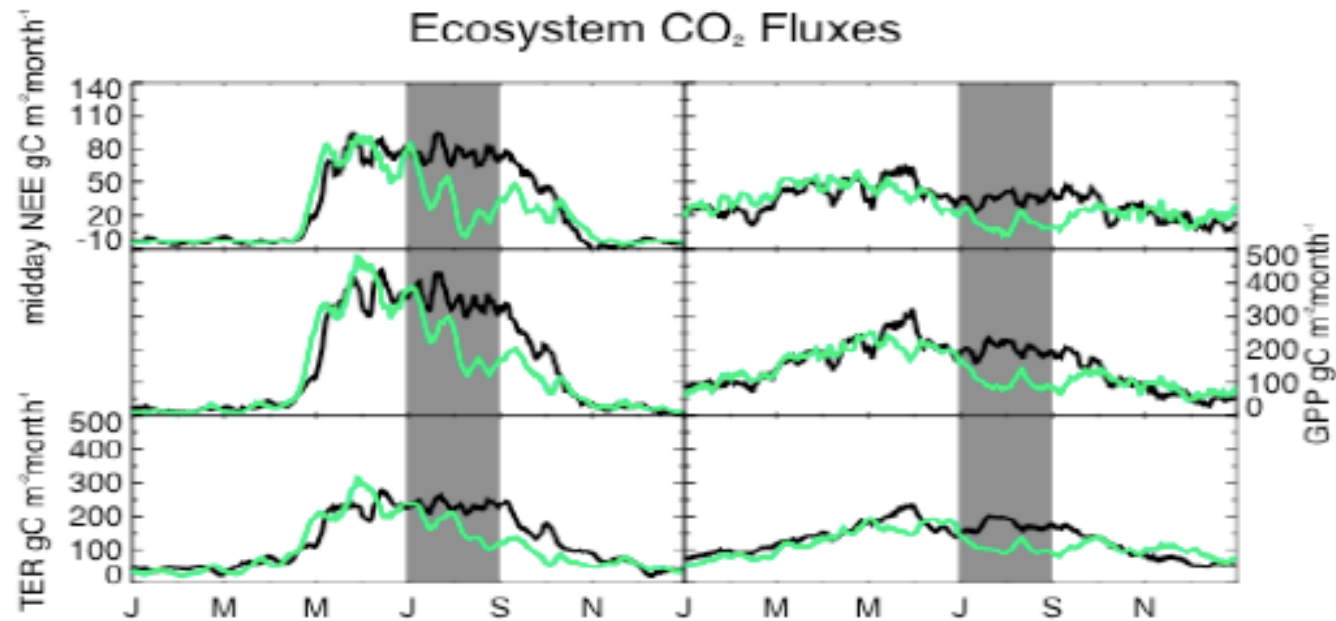
Vol 437|22 September 2005|doi:10.1038/nature03972

nature

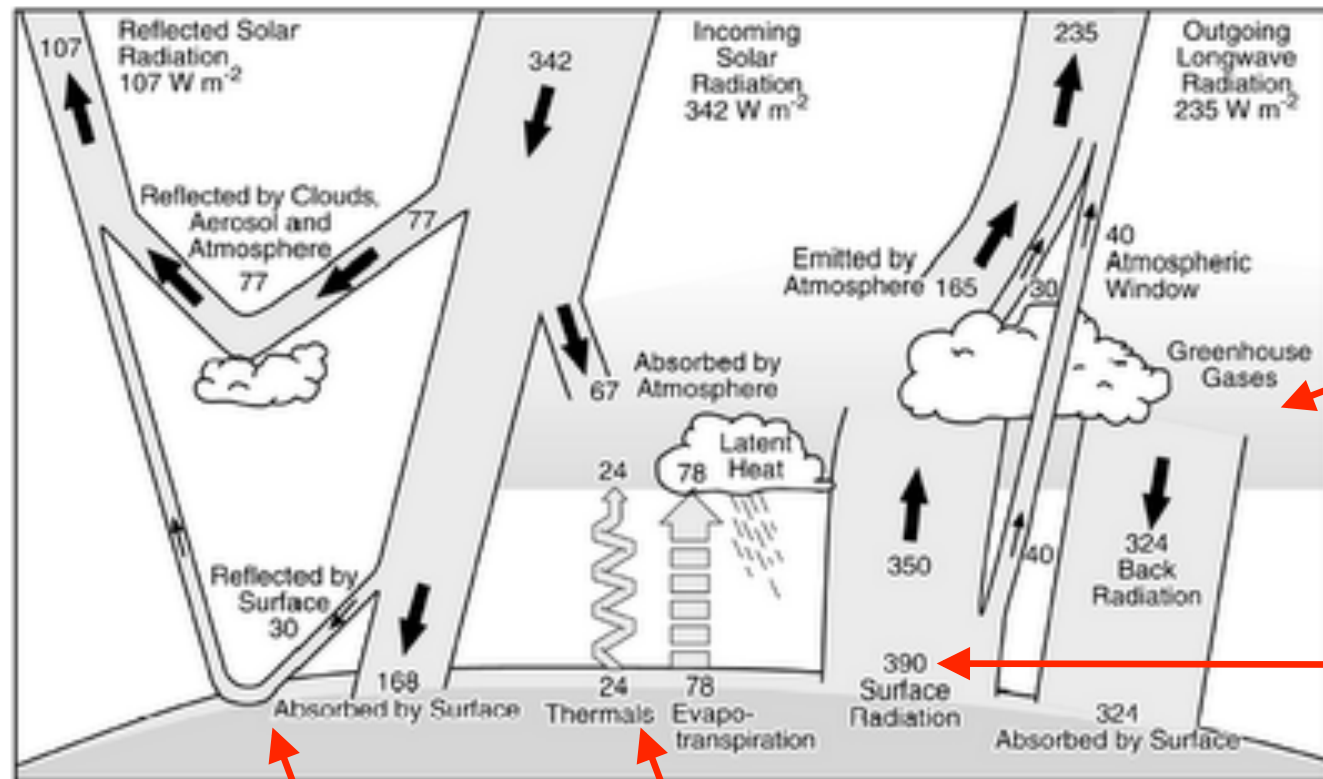
LETTERS

Europe-wide reduction in primary productivity caused by the heat and drought in 2003

Ph. Ciais¹, M. Reichstein^{2,3}, N. Viovy¹, A. Granier⁴, J. Ogée⁵, V. Allard⁶, M. Aubinet⁷, N. Buchmann⁸, Chr. Bernhofer⁹, A. Carrara¹⁰, F. Chevallier¹, N. De Noblet¹, A. D. Friend¹, P. Friedlingstein¹, T. Grünwald⁹, B. Heinesch⁷, P. Keronen¹¹, A. Knohl^{12,13}, G. Krinner¹⁴, D. Loustau⁵, G. Manca^{2†}, G. Matteucci^{15†}, F. Miglietta¹⁶, J. M. Ourcival¹⁷, D. Papale², K. Pilegaard¹⁸, S. Rambal¹⁷, G. Seufert¹⁵, J. F. Soussana⁶, M. J. Sanz¹⁰, E. D. Schulze¹², T. Vesala¹¹ & R. Valentini²



The Earth energy Balance



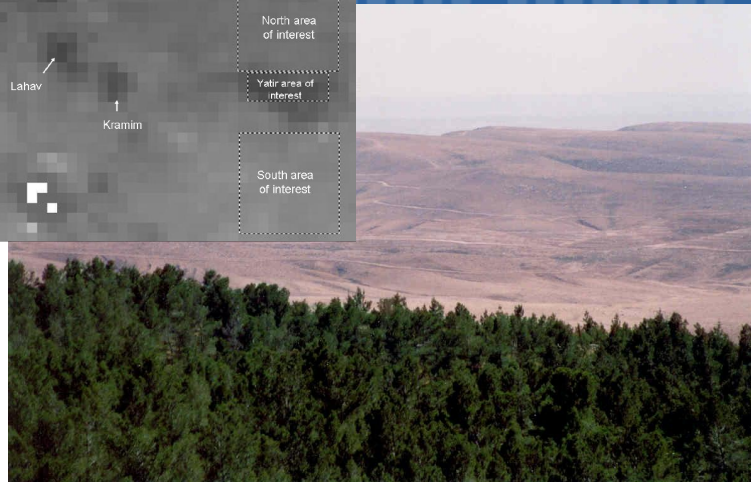
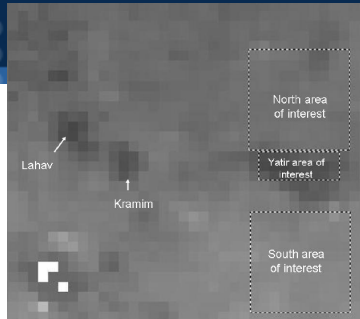
CO_2 uptake lowers
downward thermal
radiation=Cooling

Thermal radiations
temperature-
dependent

Vegetation
reduces albedo

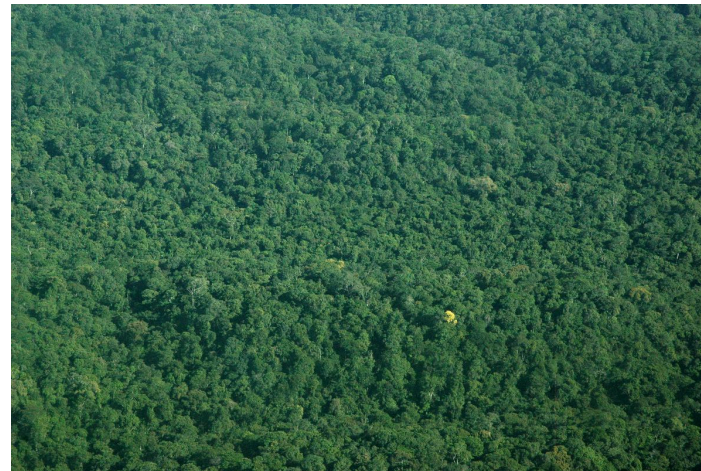
ET cools surface

Bowen ratio (H/LE) < 3 for vegetation



Albedo effect: High radiation,
low cloud, reflective background...

Water-cooled canopy... →



Air-cooled canopy...
(‘convective effect’; low r_a)



*Cooling relative to non-forest area:
~5C annual mean ←
Up to ~30C in summer mid-day

Forest/non-forest characteristics

Variable	Forest	Shrubland	
Global radiation (E_g , Wm^{-2})	238	238	
Albedo (unit-less)	0.11	0.21	$d=+23.8 \text{ Wm}^{-2}$
Net solar radiation (S_n , Wm^{-2})	212	188	
Net longwave radiation (L_n , Wm^{-2})	-96	-121	$d=+25 \text{ Wm}^{-2}$
Net radiation ($R_n=S_n+L_n$, Wm^{-2})	115	67	$d=+48 \text{ Wm}^{-2}$
Skin temperature [$^{\circ}\text{C}$]	19	24*	

Radiative forcing (RF) as a metric for the tradeoffs*:

1) Cooling effect: ~2.2 t ha⁻¹ a⁻¹ since 1965=>~27 w m⁻²

3) Warming effect: Increased surface radiation load (24 S + 25 L) ~49 w m⁻²

$$RF_{S+L}(y) = A(y)(\delta\alpha \cdot Eg + \delta L) / ES$$

3) Balance: requires ~30 more years... (Worst case scenario!)

$$y = \left(\frac{RF_{\text{surface}} \cdot \bar{\kappa} \cdot C_0}{A_E \cdot NEE \cdot \eta \cdot \zeta} \right)$$

*Radiative forcing is not climate or temperature

A=area; α =albedo; E_g =global irradiance; L=LWR; η =CO₂ RF efficiency; C=CO₂ emitted/absorbed;
 κ =convert KgC to ppmv; ζ =airborne fraction

Bottom line:

***Ecosystem adjustments help maintain high productivity in the face of warming/drying (phenology, CUE, WUE, NUE).**

***But, surface energy aspects complicate pictures and must also be considered...**

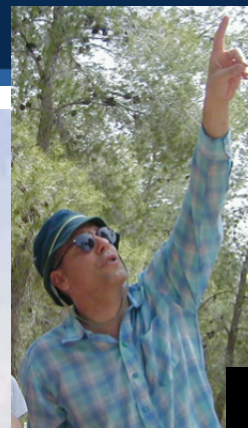
***Similarly, desertification also involves “complicating” factors of surface energy vs. carbon...**



Back in the lab



Tree rings



Energy budget



From NEE to GPP



Physiology



Ecophysiology



Stable isotopes



Where is the water



Peace of mind



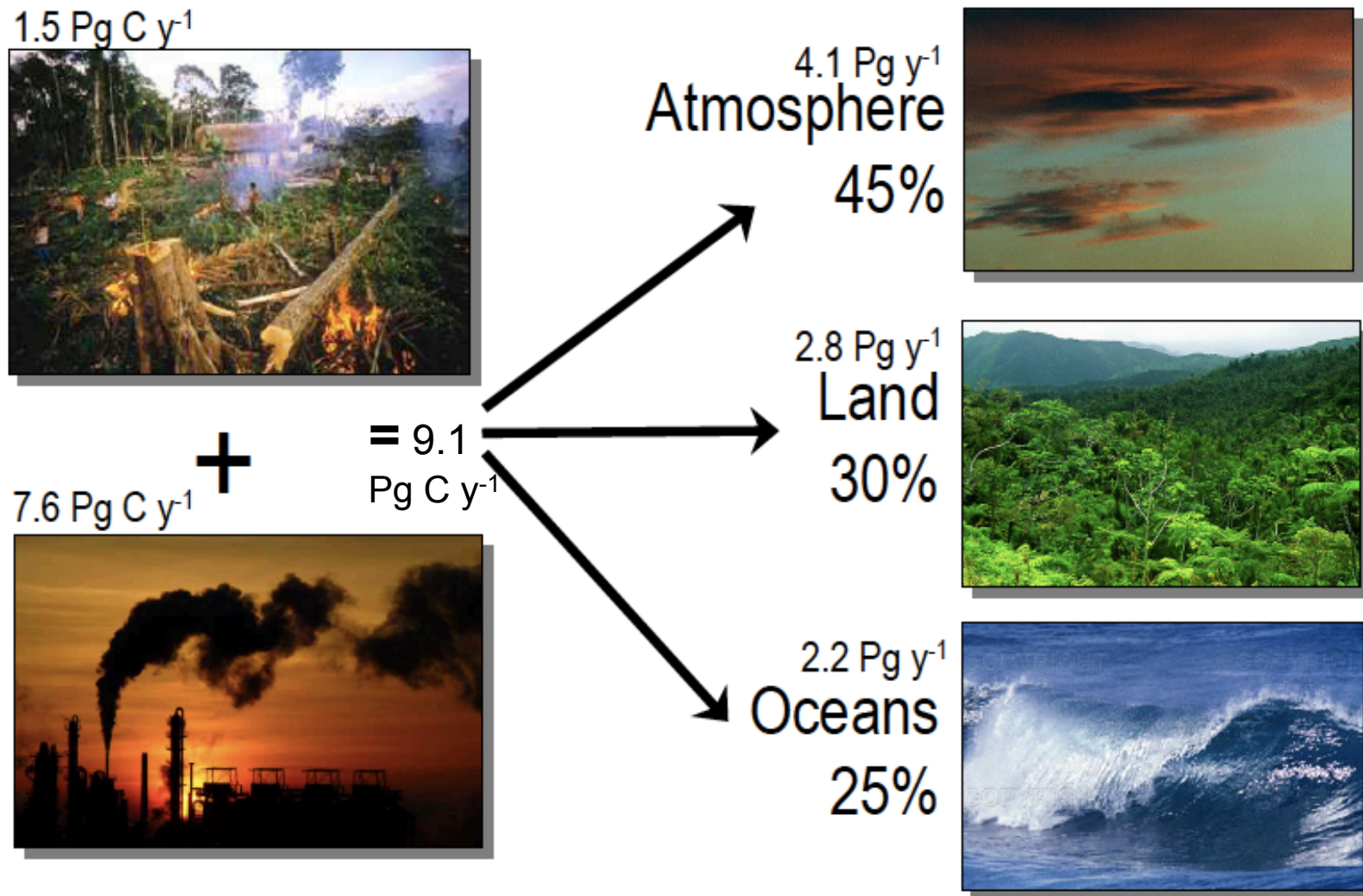
N cycle



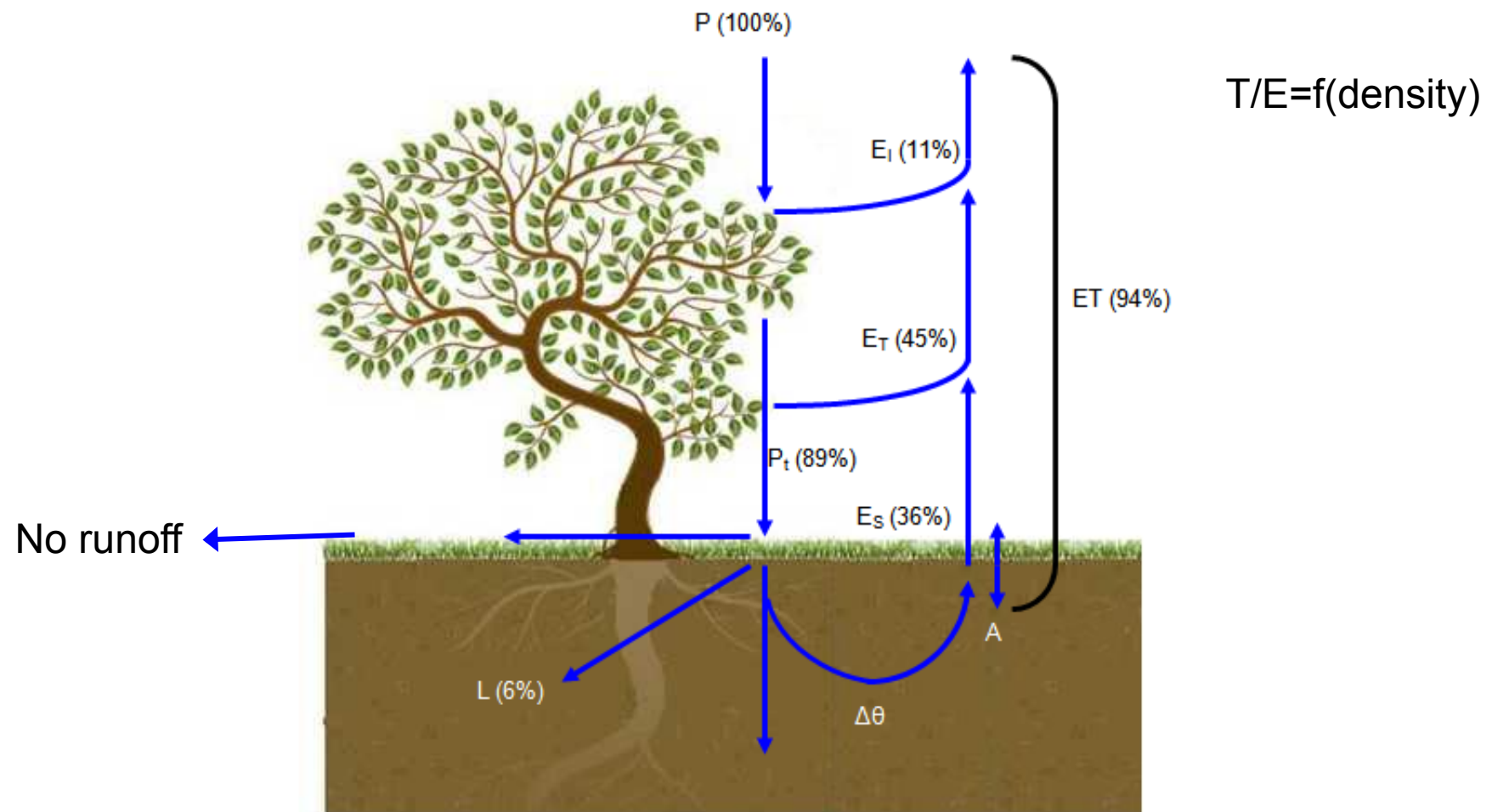
Data processing



Below ground...



Yatir's Ecohydrology



The desertification perspective

1) 35 years desertification ~5.8 Mha y⁻¹

2) Albedo + LWR suppression => RF ~0.145 W m⁻²(cooling)

$$RF_{S+L}(y) = A(y)(\delta\alpha \cdot Eg + \delta L) / ES$$

1) C release, 2.2 kg m⁻² => RF ~0.006 W m⁻² (warming)

$$RF_{CO_2}(y) = \eta \cdot Ln\left(1 + \frac{C(y)}{C_o}\right); C(y) = A(y)C\kappa / \xi$$

2) 44 ppm CO2 same period => ~0.62 W m⁻² (warming).

**3) (0.145-0.06)/0.62=~20% moderation, same period.
(similar to 'airborne fraction' moderation)**

Estimates based on: Lal 2004, Betts 2000, Myhre et al. 1998, this work

A=area; α =albedo; E_g =global irradiance; L=LWR; η =CO2 RF efficiency; C=CO2 emitted/absorbed;
 κ =convert KgC to ppmv; ξ =airborne fraction

Unique radiation budget of semi-arid forest

